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## Centroid moment tensor solutions for intermediate-depth earthquakes of the WWSSN–HGLP era (1962–1975)

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### Abstract

Centroid moment tensor solutions are presented for 76 intermediate-depth earthquakes (with reported depths between 130 and 300 km) covering the years 1962–1975. These solutions are obtained by applying the algorithm used for modern events to restricted datasets of analog (WWSSN) and digital (HGLP) seismograms. © 2001 Elsevier Science B.V. All rights reserved.

**Keywords:** Centroid moment tensor; Intermediate-depth; Earthquakes; WWSSN; HGLP

### 1. Introduction

We present a catalogue of 76 centroid moment tensor solutions for intermediate-depth earthquakes, covering the years 1962–1975, obtained by applying the inversion algorithm routinely used in the Harvard CMT project to analog records from the world-wide standardized seismograph network (WWSSN; 1962–1974), and to digital high-gain long period (HGLP) records, as well as a few early international deployment of accelerometers (IDA) records, for the year 1975. This work follows in the steps of our similar catalogue for deep earthquakes (Huang et al., 1997). The catalogue is believed to be complete for moments  $M_0 \geq 10^{26}$  dyn-cm, and more than doubles the population of reliable CMT solutions for these moment and depth windows.

We refer to Dziewonski et al. (2000) for the most recent update of the Harvard CMT catalogue, and to

Dziewonski et al. (1999) for a complete set of references to the other CMT reports published by the Harvard group over the past 17 years. In a recent contribution, Ekström and Nettles (1997) have given a modern calibration of the HGLP instruments, and provided 108 new CMT solutions, essentially extending the CMT catalogue to include the year 1976.

A full discussion of the geophysical implications of the results of this experiment will be given elsewhere. We simply present, here, a brief outline of the operational procedure used to build the WWSSN–HGLP catalogue.

### 2. Rationale

Our continued research effort is aimed at alleviating any possible undersampling of the present-day CMT catalogue, which covers only 24 years. In Huang et al. (1994), we demonstrated the possibility of using a limited number of narrow-band records (such as WWSSN or HGLP seismograms) to invert for the moment tensors of deep earthquakes, and showed in particular

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Table 1 (Continued)

No.	Centroid Parameters								Principal Axes								Best Double Couple						
	D	M	Y	Time h m sec	Latitude $\delta_{A_0}$	Longitude $\delta_{E_0}$	Depth $d_{h_0}$	Scale Factor $10^6 s$	$\sigma$	$\delta$	$\xi$	$\sigma$	$\delta$	$\xi$	$M_0$	$\varphi_s$	$\theta$	$\lambda$	$\varphi_s$	$\theta$	$\lambda$		
51	22	3	1972	10 27	50.8±0.6	8.7	45.77±0.6	-0.28	154.16±0.9	0.56	131.3±2.2	-3.7	6.5	26	2.09	29 110	-0.15	13	207	1.94	57	318	
52	22	5	1972	20 46	22.2±1.0	7.1	-17.48±1.1	0.28	-174.88±0.8	0.17	220.8±4.1	-12.8	7.3	26	4.83	9	25	-0.25	35	121	4.58	54	283
53	5	9	1972	17 18	26.9±0.9	-2.6	1.90	-	128.20	-0.17	141.0	-1.7	3.2	25	2.90	79	328	0.90	8	191	-3.79	8	180
54	5	1	1973	13 54	3.8±0.4	1.7	-35.94±0.6	0.08	175.29±0.6	0.06	161.5±1.5	14.5	4.3	26	1.03	46	22	-0.17	31	150	-0.86	28	259
55	3	4	1973	13 54	2.7±0.9	-0.35	4.36±0.9	-0.35	-75.60±0.6	0.07	145.9±1.6	-0.1	3.3	25	3.08	18	119	-0.28	7	211	-2.80	71	321
56	1	8	1973	1 31	35±0.8	4.3	-14.18±0.6	0.15	166.92±0.6	-0.37	210.6±1.3	8.6	7.6	26	3.89	79	175	-0.22	11	338	-3.37	3	69
57	30	8	1973	18 26	5.1±1.0	-0.73	7.97±1.1	0.73	-72.60±0.6	0.25	173.5±1.7	-5.5	3.5	25	3.43	44	112	-0.15	24	356	3.25	36	25
58	11	9	1973	23 18	54.6±1.1	4.2	25.44±0.9	-0.21	124.61±0.8	0.03	147.7±1.5	9.7	3.0	25	3.18	18	65	1.07	28	165	3.25	56	305
59	19	12	1973	12 56	5.1±0.6	-0.11	-20.71±0.6	-0.21	-176.28±2.0	0.04	227.9±2.0	36.9	5.1	26	1.15	42	116	-0.19	3	23	0.96	48	290
60	11	3	1974	11 37	32.7±1.0	1.1	45.52±0.7	0.21	153.11±1.6	-0.05	145.8±3.2	-8.2	2.7	25	1.47	31	267	0.46	22	11	1.93	51	130
61	17	5	1974	20 65	14.3±0.7	2.2	-6.41±0.8	0.14	106.34±0.6	-0.43	131.6±3.4	-19.4	3.6	25	3.87	20	17	0.72	55	138	-4.89	28	276
62	4	6	1974	4 14	22.9±0.7	4.2	-15.84±0.5	0.05	-174.58±1.0	0.45	233.8±1.7	27.8	5.4	26	1.30	39	121	0.20	4	28	1.50	50	293
63	30	7	1974	5 12	49.5±0.3	9.1	-36.23±0.4	-0.19	70.36±0.4	-0.40	214.2±1.6	5.2	5.2	26	4.09	58	15	-0.11	3	111	3.97	32	203
64	7	10	1974	21 53	12.2±0.7	-4.2	-58.13±0.7	-0.05	-27.78±1.0	-0.32	272.9±2.0	-13.1	3.0	25	2.71	55	220	0.23	19	342	2.94	25	81
65	8	11	1974	21 23	22.2±0.5	4.2	-42.45±0.5	-0.08	141.69±1.4	-0.06	131.6±4.2	-3.4	3.5	25	3.52	28	324	-0.37	3	355	3.15	61	151
66	5	12	1974	11 57	32.9±2.5	1.8	-8.31±2.1	-0.66	-74.33±1.1	0.13	164.0±3.1	8.0	2.0	26	1.07	6	86	-0.09	5	356	-0.98	52	225
67	15	1	1975	20 30	1.1±0.9	7.2	-7.98	-	112.30	12.9±3.7	12.9±3.7	1.9	2.5	24	2.28	53	18	-0.14	30	238	2.14	20	136
68*	17	1	1975	9 30	45.1±0.6	2.8	-17.87±0.7	0.04	-174.33±0.8	0.25	129.9±1.9	-23.1	2.5	25	5.25	4	111	-0.37	26	18	4.88	26	170
69*	25	3	1975	6 41	35.7±0.3	2.2	-27.96±0.2	0.00	-66.57±0.3	0.09	172.0±1.1	-6.0	2.7	25	2.64	3	248	-0.44	22	339	-2.20	68	151
70*	9	4	1975	6 26	28.0±0.2	5.8	-4.64±0.1	0.00	152.78±0.2	0.09	104.2±1.2	-28.8	3.9	25	8.86	50	211	0.69	0	121	9.35	40	31
71*	8	7	1975	12 4	48.4±0.3	6.0	-21.24±0.2	-0.25	94.30±0.3	-0.40	95.7±1.0	-61.3	8.3	26	3.82	50	82	-0.53	39	276	3.29	7	190
72*	10	8	1975	10 25	55.6±0.2	9.3	-22.95±0.2	-0.30	-66.85±0.2	-0.15	225.5±1.3	59.5	3.6	26	6.92	23	63	0.96	1	293	6.98	121	62
73*	23	8	1975	13 51	28.2±0.3	4.1	54.58±0.7	0.28	160.20±0.4	0.15	137.3±1.0	-3.7	2.7	25	1.13	39	305	0.16	8	298	1.27	50	109
74	30	9	1975	3 51	7.6±0.5	8.3	-9.20±0.4	0.35	-74.58±0.6	0.06	155.0±2.2	2.5	2.4	24	4.01	7	389	0.80	41	195	-3.21	48	351
75	14	10	1975	14 53	12.5±0.6	5.7	-7.01±0.7	0.05	128.95±0.6	-0.07	182.9±2.4	15.9	2.5	24	2.18	23	275	0.26	19	133	-2.43	58	138
76*	28	12	1975	15 24	57.8±0.2	7.0	-8.09±0.1	-0.11	115.00±0.2	-0.07	198.5±0.8	2.5	2.7	25	2.33	53	28	0.35	6	291	-12.68	36	198

<sup>a</sup> For explanation of headings see Dziewonski et al. (1987).



that under favorable conditions, reliable moment tensor solutions could be derived from as few records as two components at a single station. This approach was successfully applied to 104 deep-focus earthquakes ( $h \geq 300$  km) for the years 1962–1976 (Huang et al., 1997), and later to a dataset of 35 older deep events, reaching back to 1907 (Huang et al., 1998).

Huang et al. (1994) gave a detailed explanation of the feasibility of inverting depleted datasets, based on the efficient excitation of overtone surface waves by deep earthquakes, which makes up (in terms of the richness of resolving kernels) for restricted azimuthal coverage or for poor sampling in the frequency domain due to the narrow-band character of

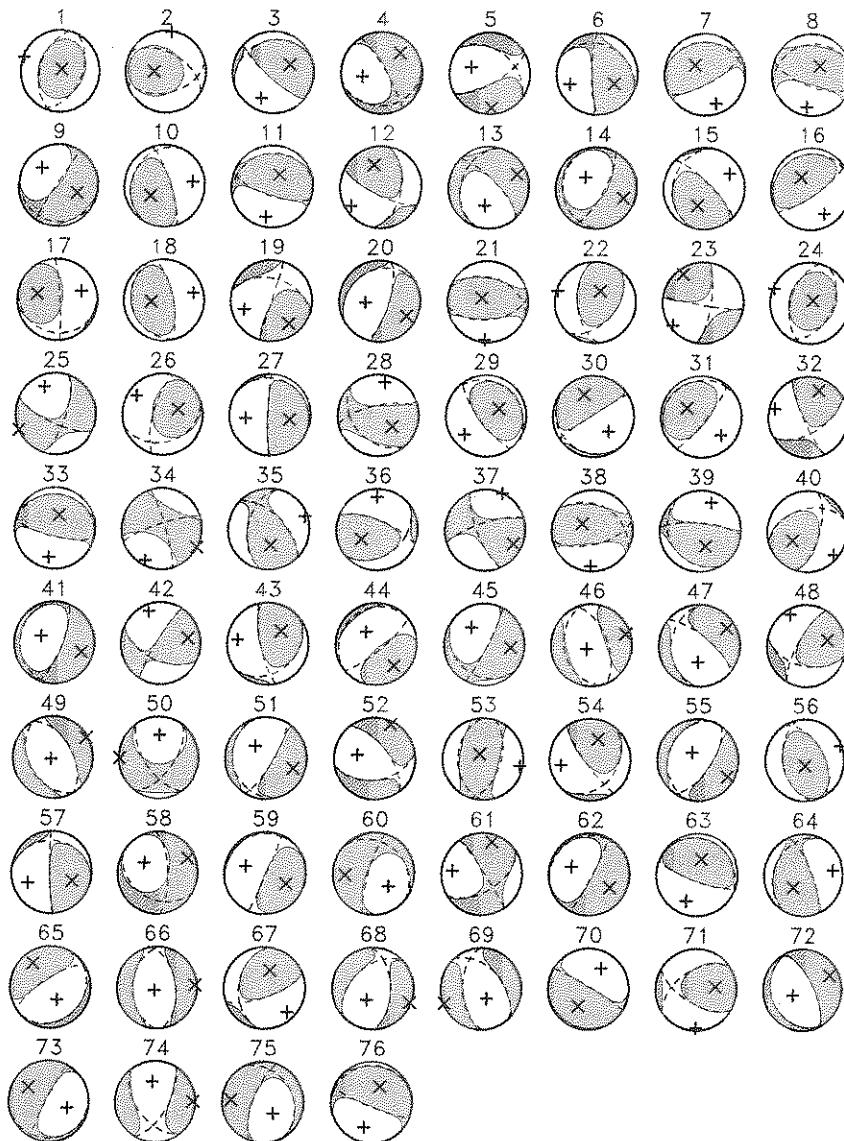


Fig. 1. Equal-area representation of the moment tensors listed in Table 2. Solid lines are the projections of the nodal surfaces of the full moment tensors; dashed lines represent the fault planes of the best double-couples, as listed in the last columns of Table 1. The compression and tension axes are shown by plus signs and crosses, respectively.

the older instruments. As the depth of the earthquake is reduced, this situation becomes less favorable, and thus, a larger number of stations must be used. In a series of systematic tests run on modern-day digital data, Huang (1996) found that the minimum number of stations necessary for a stable inversion grew from 1 at 450 to 8 at 20 km. For the intermediate depth range, these numbers would be 3 at 300 km and 5 at 130 km, further reduced to 1 and 4, respectively, if source depth is constrained in the inversion.

### 3. Selection of events

We targeted for inversion all events spanning the “WWSSN–HGLP years” (1962–1975) with a reported depth  $h$  between 130 and 300 km, and at least one reported magnitude  $M$  (most often  $m_b$ )  $\geq 5.8$ . Our experience was that smaller events could not be reliably inverted. Records from all 159 such earthquakes were visually inspected; 78 events were processed and 75 successfully inverted, among which 10 were determined from HGLP records. Unfortunately, we could not find records of sufficient quality to invert the large earthquake of 26 February 1963 in New Guinea ( $7.5^\circ\text{S}$ ;  $146.1^\circ\text{E}$ ;  $h = 156$  km); we refer the reader to Fukao and Abe (1971) for a non-CMT moment tensor solution, based on long-period Love waves. For each event to be processed, stations well distributed in azimuth were selected after inspection of the individual records, their number varying from 3 to 7 for the WWSSN (1962–1974) events, and from 3 to 12 for the HGLP-IDA (1975) events. In the case of the WWSSN records, a processing window consisting of the generalized body waves (P group, S group, and the mantle reverberations such as PS, SS, etc.) was isolated, lasting from 2 min before the P arrival to 2 min after the arrival of fundamental Love waves. Records were digitized and equalized to a common sampling  $\delta t = 1$  s, identical to that used on long-period channels of present-day digital networks. In the case of the HGLP records, we followed the procedure described by Ekström and Nettles (1997). The inversion proceeded by using exactly the same algorithm as utilized in the routine CMT determination of Dziewonski et al. (1981). Tables 1 and 2 and Fig. 1 present our dataset in the same format as used throughout the

quarterly reports and in Huang et al. (1997). Events for which mantle waves were included in the inversion (Dziewonski and Woodhouse, 1983) are identified by a star next to their number in Column 1. The format of the tables is described in detail in Dziewonski et al. (1987), to which the reader is referred. In the case of a few small earthquakes, a blank entry for the precision of the coordinates  $\lambda, \phi$  of the centroid indicates that the epicenter was fixed during the inversion. In two instances (Events 15 and 53; identified by a blank entry for  $\delta h_0$ ), we had also to constrain the depth during the inversion. With the goal of eventually merging the present catalogue with Huang et al. (1997), we keep the solution for Event 5, which converged to 361 km, and we incorporate Event 38, previously inverted to 279 km by Huang et al. (1997), but dropped from their catalogue. This brings the total number of solutions in the present dataset to 76. We also kept in the catalogue those events (numbers 1, 24, 49, 61, 65, 70 and 71) which converged to depths shallower than 130 km.

### 4. Conclusion

The total moment release for the 76 events in the catalogue is  $1.20 \times 10^{28}$  dyn-cm. We estimate that the catalogue is complete for  $M_0 \geq 10^{26}$  dyn-cm on the basis of frequency–moment statistics. This threshold is also supported by the observation that the number of solutions above the threshold (32) is comparable to that of available solutions (59) in the main CMT catalogue (1976–1999) for the relevant ranges of depth and moment, once prorated for a common sampling duration.

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